



On Extreme Winds at L-Band with the SMAP Synthetic Aperture Radar

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Abstract: In this letter, we discuss some observations of the Soil Moisture Active Passive (SMAP) mission's high-resolution synthetic aperture radar (SAR) for extreme winds and tropical cyclones. We find that the L-band cross-polarized backscatter is far more sensitive to wind speed at extreme winds than the co-polarized backscatter and it is essential to observations of extreme winds with L-band SAR. We introduce a cyclone wind speed retrieval algorithm and apply it to the limited SMAP SAR dataset of cyclones. We show that the SMAP SAR instrument is capable of measuring extreme winds up to the category 5 (70 m/s) wind speed regime providing unique capabilities as compared to traditional scatterometers with C and Ku-band radars.

Keywords: SMAP; radar; SAR; ocean winds; ocean vector winds; extreme winds; cyclone; hurricane

1. Introduction

Soil Moisture Active Passive (SMAP) is a mission designed to observe the soil moisture over land using a combined active/passive L-band system [1]. SMAP has three types of radar observations: a full-aperture measurement with approximately 36 km resolution, a range-sliced measurement with 5×36 km resolution, and a synthetic aperture radar (SAR) mode for selected portions of the Earth with a 1×1 km resolution. The SAR data is generally only available over land, however, a small amount was collected over the ocean for the descending portion of the orbit, only on the fore look, and only within 1000 km of coast. These constraints severely limited the quantity of SAR mode data available over the ocean and were driven by downlink capacity considerations. SMAP is groundbreaking in many ways: for example it is the first scanning SAR instrument and it is the first active/passive L-band instrument that is conically-scanning and thus able to estimate Ocean Vector Winds (OVW).

Unfortunately the SMAP radar failed in early July 2015, after having provided only two months of data. Though short, this dataset is sufficient to explore the vast potential of this type of system for ocean remote sensing of extreme wind events.

2. L-Band Remote Sensing of Ocean Surface

L-band active/passive remote sensing of the ocean has been proven to be a reliable means for retrieval of sea surface salinity, ocean wind speed, and ocean wind direction. Aquarius was the first combined active/passive L-band sensor flown in space and developed the field of L-band active/passive remote sensing of the ocean surface [2]. Geophysical Model Functions (GMFs) have been generated for Aquarius [3,4] and are directly applicable to SMAP after accounting for slight changes to incidence angle and resolution. Aquarius enabled combined estimates of Sea Surface

Salinity (SSS) as well as ocean wind speeds [5], and we have directly modified these algorithms as well as QuikSCAT algorithms [6–8] to the retrieval of Ocean Vector Winds (OVW), SSS, and radiometer-only extreme winds with SMAP in previous work [9–12]. In that work we show that the active/passive low resolution data, which is available for all of the time for SMAP before the radar failed, enables retrieval of OVW and SSS to accuracies approaching RapidScat and QuikSCAT for OVW and nearing but not quite as good as Aquarius for SSS. Additionally there has been some work using the SMAP SAR for retrieval of ocean surface winds [13], however, we specifically explore extreme winds. While much work has been done with C-band SAR wind estimation for both normal and extreme wind speed regimes [14,15], little has been done with L-band SAR at extreme winds.

It is known that L-band has potential for remote sensing of extreme wind events due to the longer wavelength as compared to traditional scatterometery with C and Ku-band, whereas Ku-band is significantly impacted by rain [16] and both have significant saturation or even non-monotonicity of wind speed- σ_0 relationship [17–20] at high wind speeds for the co-polarizations. L-band is not significantly affected by attenuation due to rain, and in this study we will show that the cross-polarization in particular remains sensitive to wind at the wind speeds of tropical cyclones/hurricanes, even up to category 5/70 m/s wind speeds.

3. Datasets

3.1. SMAP Data

We use the version 3 R13080 version of the SMAP L1C SAR data, available from the Alaska Satellite Facility. The L1C data were posted on a 1km grid and data was only downlinked for regions within 1000 km of coastlines, only on the descending pass, and only for the fore portion of the conical antenna scan.

3.2. Model Function

We use the GMFs derived in [3] for Aquarius for SMAP data processing. In Figure 1 we show the GMF for HH on left, VV in middle, and HV on right. Note the very large directional modulation of the L-band σ_0 for co-polarizations while the cross-polarization has much less modulation. The large directional modulation for co-polarization can cause larger speed errors due to errors in wind direction while the much smaller modulation for cross-polarization may have less errors in retrieved wind speed due to wind direction.



Figure 1. (**left**) Geophysical Model Function in dB for HH polarization, as a function of relative azimuth angle with different wind speeds as different lines, (**middle**) same for VV polarization, (**right**) same for HV polarization. Note the large wind direction modulation for co-polarizations, while the cross-polarization has much less dependance on wind direction. Also note the change in sign of directional modulation and non-monotonicity of the GMF in wind speed as cross-wind.

Also note the change in sign of the A2 term of the GMF, which for the VV polarization causes non-monotonic relation between wind speed and σ_0 for particular azimuth angles.

4. SMAP Observations of Extreme Winds

Due to the downlink constraints of SMAP there is precious little SAR data downlinked from over the ocean, and with only two months of radar data we have found very few hits on tropical cyclones. Using tropical cyclone vitals files from the National Hurricane Center (NHC) and Joint Typhoon Warning Center (JTWC) we are able to automatically collocate and extract all SMAP observations of hurricane-force winds contained in the 2 month data record. While the time of year (May-June) was not optimal for Atlantic hurricanes, the Pacific basin was quite active during 2015 and we have found about five high-quality hits on hurricane-force winds of which we show three in Figure 2. Each one clearly shows the cyclone eye, visible most easily in the cross-polarization channel. Three of the 5 hits were from East Pacific hurricanes, however, there is no collocated data available from Stepped Frequency Microwave Radiometer (SFMR) to use for validation.

For each of the scenes shown, we extract the data over each of these hits and we oversample the L1C SAR data to a 0.01° grid in latitude and longitude, with a resolution of 0.03° using a sliding window filter. Next we retrieve a wind speed from each SAR image, using a hybrid algorithm. First we perform the combined active/passive (CAP) ocean vector wind processing introduced in [9] to generate the low-resolution ocean vector winds. Then we estimate the eye location from the HV polarized σ_0 , compute the cyclonic wind direction, and linearly combine it with the wind directions from the low resolution CAP processing to generate the cyclone prior direction. This is the wind direction that is used for the subsequent high-wind speed retrieval. Finally we retrieve the wind speed on the high-resolution grid assuming the cyclone prior wind direction using the following objective function:

$$F(spd) = \left[\frac{\sigma_{0,hh} - \sigma_{0,hh}^{m}}{\sqrt{var(\sigma_{0,hh})}}\right]^{2} + \left[\frac{\sigma_{0,vv} - \sigma_{0,vv}^{m}}{\sqrt{var(\sigma_{0,vv})}}\right]^{2} + w(anc_spd)\left[\frac{\sigma_{0,hv} - \sigma_{0,hv}^{m}}{\sqrt{var(\sigma_{0,hv})}}\right]^{2}$$

Here $\sigma_{0,i}$ is the observed SAR σ_0 for polarization *i*, $\sigma_{0,i}^m$ is the model NRCS for polarization *i*, *var* ($\sigma_{0,i}$) is the expected variance of the NRCS for polarization *i*, and the HV weight *w* varies linearly from zero to one between ancillary speeds of 15 to 20 m/s, is zero below 15 m/s, and is one above 20 m/s.

In Figure 2 we show the HH σ_0 , the HV σ_0 , and the SAR wind speed retrieval for super typhoon Noul when it was a category 5 storm, cyclone Quang as a category 4 storm, and hurricane Blanca as a category 2 storm. The excluded line down the middle of the swath is the nadir region where the SAR data processing is not possible due to geometry. In each of the HH σ_0 images we see the typical hour-glass shape which is due to the high directional modulation of the HH σ_0 , in particular the second cosine harmonic; in contrast, the HV σ_0 does not display such large directional modulation. At upwind and downwind the σ_0 for HH is significantly larger than the cross-wind directions, giving rise to the hour-glass type of shape in the images on the left side of Figure 2. The HH image also does not show a clearly identifiable eye-wall location; however, the HV σ_0 does allow us to clearly delineate the eye-wall location. The eye-wall structure in Quang is tighter and better defined than in the Noul image indicating a more tightly organized storm, whereas the eye-wall in the Blanca image is quite large, somewhat disorganized, and an interesting feature of larger σ_0 directly in the middle of the eye. Also note that the somewhat odd looking wind field in the Blanca SAR wind plot does not align with the hour-glass shape in the corresponding HH image; in fact it better aligns with gradations in the HV image indicating that these may be the real winds and not an artifact due to the GMF.

Using our SAR wind speed retrieval algorithm we find wind speeds in the eye-wall between 60 and 80 m/s for Noul agreeing reasonably well with the JTWC estimates. Estimates in Quang are slightly less, perhaps overestimating as compared to the JTWC estimates, and those for Blanca seem right in line with the NHC estimates.



(a) Super Typhoon Noul on May 9th 2200 UTC; JTWC best track indicates 71 m/s



(b) Cyclone Quang April 29th 2015 approx. 2230 UTC; JTWC best track indicates 59 m/s



(c) Hurricane Blanca June 7th 2015 approx. 1340 UTC; NHC best track indicates 44 m/s

Figure 2. (a) Typhoon Noul, (b) cyclone Quang, (c) hurricane Blanca. In each we have (left) HH σ_0 , (middle) HV σ_0 , (right) SAR wind speed. The arrows on the right plots are cyclonic prior wind direction used in speed retrieval. The VV σ_0 is not shown, however, it is qualitatively very similar to the HH σ_0 . In each we can clearly see the eye-wall in the HV, however, only in the case of Blanca is the eye-wall clearly visible in the HH σ_0 . There is a distinctive hour-glass shape in each of the HH images due to the high directional modulation for co-polarizations while the HV does not show this effect.

5. Discussion and Implications on L-Band Remote Sensing of Extreme Winds

5.1. Is the HV Signal Real or Is It Leakage from Co-Polarizations?

When calibrating a cross-polarization channel, especially over the ocean where the expected cross-polarization response is very small compared to the co-polarization response, one has to be very careful to remove various cross-talk sources or at least determine if they are driving the observed signal in the cross-polarization. We would expect two main sources of cross-talk: first, due to Faraday rotation and, second, due to leakage in the instrument itself. We do not think Faraday rotation is significant for the data considered here due to two main factors: SAR data over ocean is only collected on the descending pass (6 am local time) when the total electron content (TEC) is minimal and even then a Faraday rotation correction is performed on the SAR data using a model TEC data product. Secondly, examination of the HH and HV σ_0 in Figure 2 shows distinct differences between the HV and HH channels, in particular the clear delineation of the eye wall and the hour-glass shape in the HH not being replicated in the cross-pol. In addition the cross-pol does not show nearly as much directional modulation as the co-polarizations, thus we conclude that there is significant signal in the cross-pol that cannot be due to leakage from the co-pols.

5.2. Implications on L-Band Remote Sensing of Extreme Winds

The images of the various storms in Figure 2 would seem to indicate that the HV polarization is essential to retrieval of wind speed for these extreme wind events. In Figure 3 we plot the ratio of the HV σ_0 to the geometric average of the HH and VV σ_0 . Noul had category 5 winds and Quang was nearly as strong as Noul when observed by the SMAP SAR, however it was classified as a category 4. In comparison Blanca was a category 2 storm with a much larger eye and covered a much larger area. We cannot know the exact structure of the true winds as a function of distance from the eye, however, we can assert some basic assumptions and see what conclusions may be drawn. Firstly, the maximum speed of the cyclones will typically be in the eye wall, where we observe the maximum HV backscatter in Figure 2a,b in the middle plot of each. Secondly, we can assume that for at least a small distance from the eye that the azimuthally-averaged wind speed is monotonically decreasing as a function of distance from eye, for distances larger than the eye radius. From these two assumptions and the data shown in Figure 3 we can conclude that the HV σ_0 has an increasing sensitivity as compared to the HH or VV σ_0 as wind speeds approach and exceed the category 4 and 5 wind speed regimes. For normal wind speeds, the Aquarius GMF derived in [3] gives a cross-pol to co-pol ratio of -16.6 dB at 7.5 m/s wind speeds (see green curves in Figure 1). We conclude that this increase in sensitivity of the cross-pol channel as compared to the co-pol channels is about 4–5 dB in magnitude from the normal wind speed to category 5 wind speed regimes. This effect is consistent in both of the SMAP SAR observations of category 4 and 5 wind speeds in Noul and Quang.

There is also some evidence of possible saturation of the co-polarization σ_0 channels for these extreme winds. We see that the images of Noul and Quang (Figure 2a,b) that the cross-pol σ_0 clearly shows the eye-wall locations while the HH σ_0 does not. In both cases the co-pol channel also does not show a strong increase in σ_0 in the vicinity of the eye-wall, but rather a much weaker relative change as compared to further away. In contrast, in the images of Blanca (Figure 2c) the HH σ_0 and HV σ_0 both clearly illustrate the eye-wall location. Perhaps somewhere between the category 2 and category 4–5 scale the co-polarization σ_0 starts to lose sensitivity to changes in the wind speed. Combined, these observations demonstrate that HV will be a necessity to observe extreme winds with L-band SAR.



Figure 3. Ratio of the HV σ_0 to the geometric average of the HH and VV co-polarization σ_0 plotted as a function of distance from estimated hurricane eye location for Noul (black), Quang (red), and Blanca (blue). The cross-pol co-pol ratio seems to peak at -12 dB for extreme winds and decreases as wind speed decreases. The Aquarius GMF indicates -16.6 dB ratio at 7.5 m/s wind speed. Note that Blanca, while weaker, is spatially a much larger storm with a larger eye, has two peaks (in agreement with the wind fields), and does not reach as high of a cross-pol co-pol ratio as the other two storms (less intense).

6. Conclusions

We have found collocations of SMAP SAR data with a few tropical cyclones/hurricanes. We show that the L-band SAR signal shows strong sensitivity to extreme winds up to the category 2 wind speed regime for the co-polarizations and up to the category 5 regime for the cross-polarizations. The cross-polarization increases in sensitivity as compared to the co-polarizations as winds increase from the normal to extreme wind regimes. It is unclear if the HH polarization saturates between the category 2 and 5 wind regimes, however, the HV clearly does not. We have devised a wind speed retrieval method for the SMAP SAR data and shown it seems to give reasonable agreement with what data are available. With precious little data, it is impossible to validate these hurricane-force wind retrievals, however, the data clearly show that L-band is capable of measuring these extreme winds in a way that other Ku and C-band scatterometers cannot.

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